

Evaluation of the Influence of the Physico-Chemical Form of ^{131}I on the Thyroid Dose Estimates

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aid dose estimates that are calculated in the report are based on the assumption that ^{131}I in the radioactive cloud is only in particulate form. This appendix addresses the question of the influence of the simplifying assumption on the estimated doses that may have resulted. The alternative approach has been to make more complex calculations that take into account the behavior of various forms explicitly.

BACKGROUND AND ASSUMPTIONS

The physico-chemical forms of airborne ^{131}I can be classified

- iodine associated with particles
- molecular (I_2)
- organic (such as CH_3I).

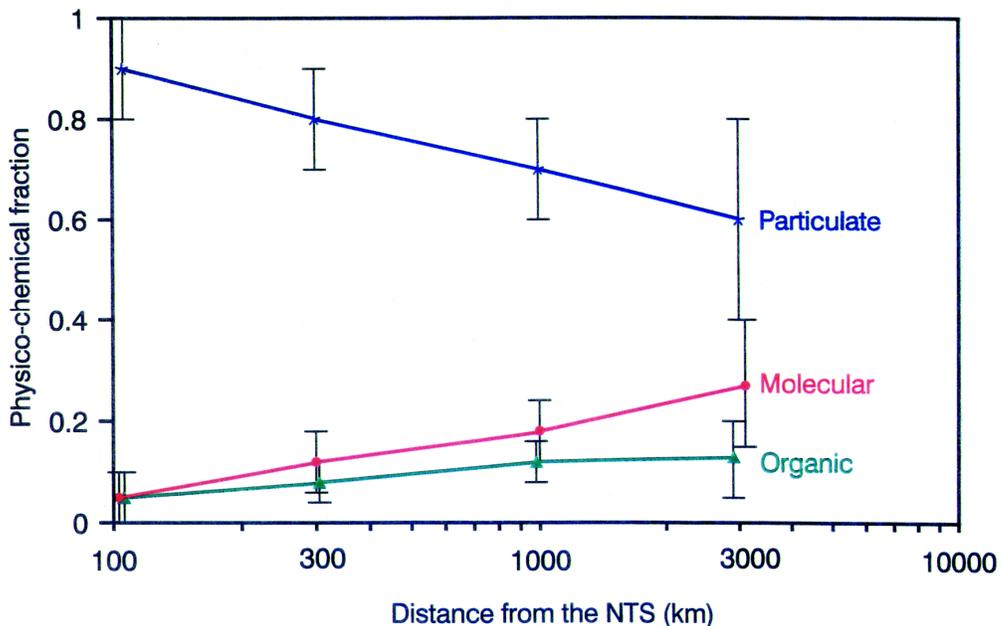
It is thought that:

Most of the ^{131}I was in particulate form close to the NTS. There was a broad range of particle sizes soon after detonation. As the radioactive cloud moved away from the NTS, the inventory of large particles was progressively depleted due to deposition. The median particle size in the cloud decreased with increasing travel time and distance from the NTS. Simultaneously, chemical reactions produced organic and molecular forms.

- The fractions of elemental iodine and organic forms were small in the initial radioactive cloud but became larger with increasing time since detonation (or at greater distances from the NTS).

On the basis of observations of fallout from Chinese weapons tests in recent years for long distances from the NTS (Voilleque 1979) and on measurements following the Babel event of December 1970 for short distances from the NTS (Pendleton et al. 1971), the relative distribution of the physico-chemical forms (PCF) considered is assumed to vary as a function of distance X from the NTS as shown in Figure A7.1 in Table A7.1.

7.1. Relative distribution of the physico-chemical forms of ^{131}I according to distance from the NTS. (The vertical bars represent the assumed ranges of uncertainty)



1. Best estimates and ranges of the fractions of ¹³¹I associated with particles, in molecular form, and inorganic form, according to distance, X, from the NTS.

| chemical form | X = 100 km | | X = 300 km | | X = 1000 km | | X = 3000 km | |
|---------------|------------------|---------|------------------|-----------|------------------|------------|------------------|----------|
| | Estimated values | | Estimated values | | Estimated values | | Estimated values | |
| | Best | Range | Best | Range | Best | Range | Best | Range |
| particle | 0.9 | 0.8-1.0 | 0.8 | 0.7-0.9 | 0.7 | 0.6-0.8 | 0.6 | 0.4-0.8 |
| molecular | 0.05 | 0-0.1 | 0.12 | 0.06-0.18 | 0.18 | 0.12- 0.24 | 0.27 | 0.15-0.4 |
| inorganic | 0.05 | 0-0.1 | 0.08 | 0.04-0.12 | 0.12 | 0.08-0.16 | 0.1 | 0.05-0.2 |

For the purposes of the uncertainty analysis, the statistical distributions of the molecular and organic form fractions are to be uniform within the specified ranges. In the calculation procedure, the fraction of ¹³¹I associated with particles is by subtraction of the sum of the values for the molecular and organic forms; its statistical distribution is observed to be approximately triangular.

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Thyroid doses arise from inhalation of ¹³¹I-contaminated air and ingestion of ¹³¹I-contaminated foodstuffs. A comparison of thyroid doses obtained when ¹³¹I is only in particulate form and when it is distributed among various physico-chemical forms is required to determine whether the dose estimates calculated under these assumptions are greatly different. Because most of the thyroid dose is proportional to the deposition density of ¹³¹I on vegetation (also called "¹³¹I vegetation deposition density" in this report), and since the relationship between the thyroid dose and ¹³¹I vegetation deposition density is independent of the chemical form of ¹³¹I, the comparison of the dose estimates is to a large extent equivalent to the comparison of the deposition densities of ¹³¹I on vegetation.

The comparison of the ¹³¹I vegetation deposition densities when ¹³¹I is only in particulate form and when ¹³¹I is distributed among various physico-chemical forms was carried out using the unit time-integrated concentration of ¹³¹I in ground-level air (IC_{air}) of 1 nCi d m⁻³ (equivalent to an average concentration of ¹³¹I in ground-level air of 1 nCi m⁻³ over 1 day).

The parameters influenced by the physico-chemical form are:

the dry deposition velocity, v_g,

the washout ratio, WR,

the mass interception factor, F*

The calculations of the vegetation deposition densities were performed for the following conditions:

- 4 distances X from the NTS: 100 km, 300 km, 1000 km, 3000 km
- dry deposition and wet deposition for eight values of daily rainfall, R, that are representative values for the precipitation index (Table A7.2).
- 3 physico-chemical forms (particles (P), molecular and organic (ORG)), the relative distributions of which are assumed to vary as a function of distance from the NTS as described above (see Figure A7.1).

The vegetation deposition density, A_p in nCi m⁻², is the product of the deposition density on the ground, DG in nCi and of the interception fraction, F:

$$A_p = DG \times F$$

The deposition density on the ground is:

- under dry conditions, obtained as the product of the unit time-integrated concentration in ground-level air, IC_{air} and of the dry deposition velocity, v_g:

$$DG_{dry} = IC_{air} \times v_g$$

where:

$$IC_{air} = 1 \text{ nCi d m}^{-3}$$

v_g is expressed in m d⁻¹

amount of daily rainfall, R, and the density of the air at level, AD:

$$DG_{wet} = IC_{air} \times R \times \frac{WR}{AD} \quad (A7.3)$$

λ is the ratio of the time-integrated concentrations of ¹³¹I in rain and in air, expressed in (nCi d kg⁻¹)/(nCi d kg⁻¹)⁻¹

ρ is the density of air at ground level = 1.2 kg m⁻³

The total deposition density on the ground, DG_{tot} is the sum of the depositions due to dry and wet processes:

$$G_{tot} = DG_{dry} + DG_{wet} = IC_{air} \times \left(V_g + \frac{(R \times WR)}{AD} \right) \quad (A7.4)$$

The interception fraction is the product of the mass interception factor, F*, and of the biomass, Y:

under dry conditions:

$$F_{dry} = F^*_{dry} \times Y \quad (A7.5)$$

for deposition with rain:

$$F_{wet} = F^*_{wet} \times Y \quad (A7.6)$$

F*_{dry} and F*_{wet} are the mass interception factors for dry and wet respectively, expressed in m² kg⁻¹ (dry weight of vegetation).

The estimates of the vegetation deposition densities given physico-chemical forms, daily rainfalls, and distance to the NTS, are derived from equations A7.1 to A7.6:

- under dry conditions:

$$A_{p,dry}(X, R, PCF) = IC_{air} \times v_g(X, PCF) \times F^*_{dry}(X, PCF) \times Y$$

- deposition with rain:

$$A_{p,wet}(X, R, PCF) = IC_{air} \times R \times WR(X, R, PCF) \times F^*_{wet}(X, PCF) \times Y$$

- for the total deposition:

$$A_{p,tot}(X, R, PCF) = A_{p,dry}(X, R, PCF) + A_{p,wet}(X, R, PCF)$$

2. Daily rainfalls associated with each precipitation index.

| Weather conditions | Precipitation, indices and amounts | | Representative daily rainfall (mm, L/m ² or Kg/m ²) |
|--------------------|------------------------------------|---|--|
| | Index | Daily rainfall range (mm, L/m ² or Kg/m ²) | |
| Dry | 1 | 0 | 0 |
| Wet | 2 | > 0 - 0.25 | 0.15 |
| Wet | 3 | > 0.25 - 0.76 | 0.5 |
| Wet | 4 | > 0.76 - 2.5 | 1.5 |
| Wet | 5 | > 2.5 - 7.6 | 5 |
| Wet | 6 | > 7.6 - 25 | 15 |
| Wet | 7 | > 25 - 76 | 50 |
| Wet | 8 | > 76 - 127 | 100 |
| Wet | 9 | > 127 | 150 |

re-integrated concentration in air of ^{131}I with the mix of physico-chemical forms, $A_{p,tot}(X,R,MIX)$, is by weighting the results calculated for each physico-chemical form according to the fraction $FR(PCF)$ of ^{131}I in each chemical form:

$$\begin{aligned}
 MIX &= A_{p,tot}(X,R,P) + A_{p,tot}(X,R,M) + A_{p,tot}(X,R,ORG) \\
 &= IC_{air} \times Y \times \left[\left(FR(P) \times v_g(X,P) \times F_{dry}^*(X,P) \right) \right. \\
 &\quad + \left(FR(M) \times v_g(X,M) \times F_{dry}^*(X,M) \right) \\
 &\quad + \left(FR(ORG) \times v_g(X,ORG) \times F_{dry}^*(X,ORG) \right) \\
 &\quad + \left(\frac{R}{AD} \right) \times \left(FR(P) \times WR(X,R,P) \times F_{wet}^*(X,P) \right) \\
 &\quad + \left(FR(M) \times WR(X,R,M) \times F_{wet}^*(X,M) \right) \\
 &\quad \left. + \left(FR(ORG) \times WR(X,R,ORG) \times F_{wet}^*(X,ORG) \right) \right]
 \end{aligned}
 \tag{A7.10}$$

VALUES OF THE PARAMETER VALUES

Dry Deposition Velocity

Deposition velocity can be experimentally defined as the ratio of the activity deposited per unit area of ground and of integrated concentration in ground-level air in the absence of precipitation.

Fraction associated with particles

Deposition velocity for particles depends on the particle size. Large particles (greater than 20 μm in diameter) are removed from the air mainly by sedimentation; smaller particles are removed from the air by impaction and turbulent diffusion. Information on particle sizes in ground-level air near the fallout from atmospheric tests is available mostly for the Johnston Atoll series (Cederwall et al. 1990). Activity median aerodynamic diameters (AMADs) and geometric standard deviations (GSDs) for particle-size distributions observed in St. Johns Island, situated approximately 200 km from the NTS are given in Table A7.3 for the tests Annie and Harry. The results for Annie are much more typical of the tests sampled in the Johnston Atoll series than those for Harry (Cederwall et al. 1990). Particle size distributions for both tests are characterized by large GSD values and AMAD values that tended to increase with time.

Assuming that the upper and lower values of the ranges of particle sizes are adequately represented by the values of the AMADs, by multiplying and dividing the median estimates by the values of the GSDs, the particle-size spectrum is from 0.15 to 3,000 μm at 200 km from the NTS.

This distribution is reflected in the estimates of dry deposition velocity near the NTS. Cederwall et al. (1990) identified several locations within 320 km of the NTS where paired values of the time-integrated concentration in air and the deposition density are available from a number of tests. The resulting estimates of v_g from 168 paired values varied over several orders of magnitude with a geometric mean of 3700 m d^{-1} and a GSD of 15. This extremely large value may be partly due to the fact that the data were not corrected according to downwind distance from the NTS or to location relative to the fallout-pattern centerline. In addition, the deposition density and the time-integrated concentration in air may have been measured in the same general area but not necessarily at the same site.

Beyond 320 km of the NTS, the average particle size is expected to be smaller, resulting in turn in smaller deposition velocities. A geometric mean of 2500 m d^{-1} with a GSD of 15 was obtained from 52 pairs of air samples and deposition densities related to the Tumbler-Snapper series (List 1953).

Dry deposition velocities at greater distances from the NTS have been derived from air concentrations and deposition measurements by Pelletier and Voilleque (1971) of fallout from tests conducted in the Pacific and in territories of the former Soviet Union. The measurements took place in Michigan from 1962 to 1964. A distinction was made between fresh fallout (defined by the presence of measurable quantities of ^{14}C in air) and old fallout (absence of ^{140}Ba in air). According to this criterion, measurements in the contiguous U.S. following nuclear testing at the NTS would have been defined as fresh fallout. The average values of v_g were computed from measurements to be 1000 m d^{-1} during periods of fresh fallout and 300 m d^{-1} during periods of old fallout (Pelletier and Voilleque 1971), showing presumably the effect of smaller particle sizes for old fallout. The GSDs associated with the results were not reported but the values of v_g during periods of fresh fallout indicated to be quite variable from week to week.

On the basis of the experimental values that have been reported, the variation with distance from the NTS of the average values of $v_g(X,P)$ for ^{131}I associated with particles is found to be relatively well modeled with a power function. The empirical function that has been adopted is:

$$v_g(X,P) = 20150 \times X^{-0.35}$$

where

$v_g(X,P)$ is in m d^{-1} , and X is in km.

The distribution of $v_g(X,P)$ is assumed to be log-triangular for the four distances considered with the spread of values decreasing with distance as shown in Table A7.4. For the distance closest to the NTS, $X = 100$ km, the distribution of v_g is assumed to have a mode equal to the best estimate and to range from 400 to 200,000 m d^{-1} . This distribution corresponds approximately to the central portion (20th to 80th percent) of a log-normal distribution with a GSD of 7. Deposition velocities of 400 m d^{-1} or less are associated with very small fallout particles which are common near the NTS. The average

sition velocities at other distances as well. The indicated modal and maximum deposition velocities with is intended to reflect the depletion of large particles cloud. The assumed variation of the modes and ranges of the dry deposition velocity for ¹³¹I associated with as a function of the distance from the NTS, is illustrated A7.2.

. Molecular fraction

deposition velocities for the molecular fraction of iodine ally been derived from field experiments in which meas have been made of the time-integrated concentration l of the activity deposited on vegetation cut at between :m above ground. The dry deposition velocities obtained ay are smaller than those derived from measurements of activity deposited on the ground since the additional n on the remaining vegetation, detritus, root mat, and ot been included. Vegetation deposition velocities for r iodine vary as a function of meteorological parameters eed, air temperature, and humidity) and of the biomass; and Sauve (1981) suggested that the vegetation deposi- ity is proportional to the biomass and to the wind speed ases by a factor of 2 for a temperature increase of 10°C rease in the relative humidity of 25%.

order to avoid confusion between the two quantities n deposition velocity and total deposition velocity) and e the influence of the biomass, Hoffman (1977) recom- that the vegetation deposition velocity should be nor- or biomass (dry weight per square metre of ground). alized vegetation deposition velocity v_D is expressed in rt in m³ kg⁻¹ d⁻¹ and is related to the dry deposition according to:

$$v_D = v_g \times F^* \tag{A7.12}$$

Zimbrick 1970; Vogt et al. 1976), Hoffman (1977) estimate a value for v_D of 0.1 m³ kg⁻¹ s⁻¹ is suitable for a generic asse ment calculation; the same value, which corresponds to 900 kg⁻¹ d⁻¹, is adopted in this report. The distribution of v_D(M) assumed to be log-normal with a GSD of 2.0, at all downw distances.

A7.4.1.3. Organic fraction

The major organic form of iodine that is found in the atmo- after a nuclear test is methyl iodide (CH₃I). Methyl iodide is that deposits on vegetation to a much smaller degree than c molecular iodine. The ratio of the vegetation deposition vel (or of the normalized vegetation deposition densities) of mo- lar and organic iodine has been reported to be approximate 100 (Nakamura and Ohmomo 1980a; Nakamura and Ohr 1980b) and 200 (Heinemann and Vogt 1980). In this repor average ratio of 150 has been used, corresponding to a nor- ized vegetation deposition velocity for organic iodine, v_D(O of 9000/150 = 60 m³ kg⁻¹ d⁻¹. This value is assumed to be in- pendent of distance from the NTS. The distribution of v_D(C is also assumed to be log-normal with a GSD of 2.0.

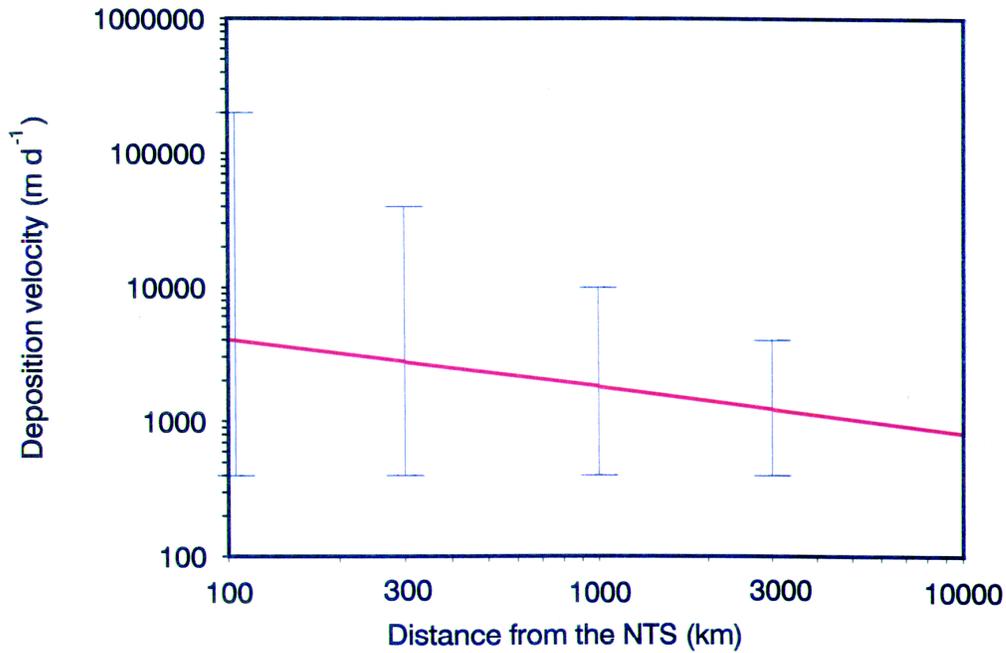
A7.4.2. Washout Ratio

The washout ratio, WR, is the ratio of the concentration of rain to that in air at ground level. The washout ratio is dim- sionless but it has a different value according to whether th concentrations are expressed per unit mass or per unit volume. this report, the concentrations are expressed per unit mass kg⁻¹).

.3. Variation of particle size with time at St. George, UT, for the tests Annie and Harry of the Upshot-Knothole series (Cederwall et al. 1990).

| Test Annie, 1320 GMT, 17 March 1953 | | | Test Harry, 1205 GMT, 19 May 1953 | | |
|-------------------------------------|------------|-----|-----------------------------------|-----------|-----|
| times,GMT | AMAD (μm) | GSD | Sample times, GMT | AMAD (μm) | GSD |
| 20 | 76 | 9 | 1310-1830 | 380 | 10 |
| 10 | 24 | 8 | 1840-2145 | 160 | 8 |
| 00 | 7.3 | 7 | 2150-0600 | 320 | 9 |
| | | | 0605-1335 | 12 | 6 |

7.2. Assumed variation of the dry deposition velocity ($m d^{-1}$) according to distance from the NTS, for ^{131}I associated with particles. (The vertical bars represent the assumed ranges of uncertainty)



4. Variation of the dry deposition velocity of ^{131}I associated with particles according to distance, X, from the NTS.

| km) | Type of distribution | $v_g(X,P)$ ($m d^{-1}$) | | |
|-------|----------------------|---------------------------|------|---------|
| | | Minimum | Mode | Maximum |
| 00 | Log-triangular | 400 | 4000 | 200,000 |
| 100 | Log-triangular | 400 | 2700 | 40,000 |
| 1000 | Log-triangular | 400 | 1800 | 10,000 |
| 10000 | Log-triangular | 400 | 1200 | 4,000 |

Washout ratio for particles has been found to increase with distance (Gatz 1977). Increasing the rainfall rate at a particular distance will lead to an increase in washout at that moment (Gatz 1978). However, washout ratios for ^{90}Sr measured for monthly storms decreased when the total amount of rain for that month increased (Krey and Toonkel 1977). A similar trend was observed when the washout ratio was plotted against the total precipitation (Krey and Toonkel 1977). Gatz (1977) observed a negative correlation with storm rainfall for washout ratios for most of the elements measured, but the trends were not statistically significant when analyzed statistically.

For the present analysis, washout ratios that are consistent with the previous analysis of scavenging factors and relative contributions of wet and dry deposition for U.S. locations have been selected (Chapters 3 and 7 and Appendix 1). Two reference values were selected: (a) 13000 for a daily rainfall amount of 1 mm for large particles ($X = 100$ km); and (b) 3000 for a daily rainfall amount of 1 mm and for small particles ($X = 3000$ m). Washout ratios for other situations were derived from the reference values.

For a given distance, X , as a function of daily rainfall, R :

$$WR(X, R, P) = WR(X, 1 \text{ mm d}^{-1}, P) \times R^{0.7} \quad (\text{A7.13})$$

For a given daily rainfall, R , as a function of distance X :

$$WR(X, R, P) = WR(100 \text{ km}, R, P) \times \left(\frac{X}{100}\right)^{-0.43} \quad (\text{A7.14})$$

The washout ratios obtained in this way are shown in Figure A7.3 for ^{131}I associated with particles and in Figure A7.4 for molecular and organic forms, for a distance from the NTS of 100 km and for the range of daily rainfall considered in this report. The higher washout ratios near the NTS reflect the presence of larger particles in the radioactive cloud. The ratios decrease more sharply with daily rainfall than the results reported by Krey and Toonkel (1977) for monthly averages or by Gatz (1978).

4.3. Molecular fraction

Values of $WR(R, M)$ for ^{131}I in molecular form are independent of distance, X , from NTS. However, they also decrease with rainfall amount (Coleman and Postma 1970). A reference value of 6000 was selected for a daily rainfall of 1 mm, based on the partition coefficient estimate of Coleman (1970) and typical airborne iodine concentrations. For other rainfall amounts were derived from the reference value using the same decrease with daily rainfall as for the

The values of $WR(R, \text{ORG})$ for ^{131}I in organic form also are independent of the distance, X , from NTS. Based on the difference in partition coefficients (Postma 1970), the reference value for a daily rainfall of 1 mm is taken to be 10. The variation function of daily rainfall has been assumed to be the same as that described above for particles and for elemental iodine.

4.4.2.4. Summary

Table A7.5 summarizes the best estimates of the washout ratios expressed in $(\text{nCi kg}^{-1})(\text{nCi kg}^{-1})^{-1}$, for the three species of ^{131}I and the four distances considered in this report. Table A7.5 includes the range of washout ratio values that is expected in each case. At a given distance from the NTS, for a given particle size and chemical form and rainfall category, the lowest value (minimum) of the washout ratio is assumed to be equal to the best estimate in the next category of higher rainfall, while the highest value (maximum) is assumed to be equal to the best estimate in the adjacent category of lower rainfall. In rainfall category 1 ($0 < R < 0.25$ mm), the maximum value was taken to be twice the mode. In rainfall category 9 ($R > 127$ mm), the minimum value was taken to be the mode divided by 1.5. The distribution of the wash-out ratios in each category is assumed to be triangular, the mode being equal to the best estimate.

4.4.3. Mass Interception Factor, F^*

The mass interception factor represents the quotient of the radionuclide concentration in vegetation (dry weight) and ground deposition density, immediately after deposition. It is expressed in $\text{m}^2 \text{kg}^{-1}(\text{dry})$. The mass interception factor is denoted as F^*_{dry} and as F^*_{wet} for deposition under dry and wet conditions, respectively.

4.4.3.1. Fraction associated with particles

Values of F^*_{dry} and F^*_{wet} for particles are calculated using the following equations, which are discussed in Chapter 4:

- Dry deposition (F^*_{dry}):

$$F^*_{\text{dry}}(X, P) = \frac{1 - e^{-\alpha(X)Y}}{\alpha(X)}$$

with

$$\alpha(X) = 7.01 \times 10^{-4} X^{1.13}$$

and

$$Y = 0.3 \text{ kg m}^{-2}, \text{ dry weight.}$$

$2.8 \text{ m}^2 \text{ kg}^{-1}$, which is met for any distance X greater than 10 km.

Wet deposition ($R \geq 5 \text{ mm d}^{-1}$) (F_{wet}^*):

$$F_{\text{wet}}^*(R, P) = 0.9 + \left(\frac{11}{R}\right) \quad (\text{A7.17})$$

respective of the distance X from the NTS (i.e., no dependence on the particle size is considered).

Wet deposition ($2.5 \text{ mm d}^{-1} < R < 5 \text{ mm d}^{-1}$):

$$F_{\text{wet}}(R, P) = F_{\text{wet}}^*(5 \text{ mm d}^{-1}) = 3.1 \text{ m}^2 \text{ kg}^{-1} \quad (\text{A7.18})$$

respective of the distance X from the NTS (i.e., no dependence on the particle size is considered).

Wet deposition ($R < 2.5 \text{ mm d}^{-1}$):

$$F_{\text{wet}}(R, P) = F_{\text{dry}}^*(X, P) + (F_{\text{wet}}^*(R_1, P) - F_{\text{dry}}^*(X, P)) \left(\frac{R}{R_1}\right) \quad (\text{A7.19})$$

$= 2.5 \text{ mm d}^{-1}$.

Molecular fraction

F_{dry}^* for the molecular fraction are not needed as the values of the vegetation deposition density for that physico-chemical form make use of the normalized vegetation deposition density which is the product $v_g \times F_{\text{dry}}^*$. In case of precipitation, values of F_{wet}^* for ^{131}I in molecular form are assumed to be smaller than the values obtained for particles far away from the NTS (Hoffman et al. 1989).

Organic fraction

F_{dry}^* for the organic fraction are not needed for the reason as for the molecular fraction. Values of F_{wet}^* for ^{131}I in organic form are assumed to be equal to the values used for molecular form.

Summary

Table 6 summarizes the estimated values of the mass interception factor. Hoffman and Baes (1979) found that the values of the mass interception factor for dry conditions are log-normally distributed with a geometric standard deviation of 1.5. In this report, the distribution of the values of the mass interception factor is assumed to be log-normal with a geometric standard deviation of 1.5, for all physico-chemical forms, daily rain-distances from the NTS.

associated with particles is illustrated:

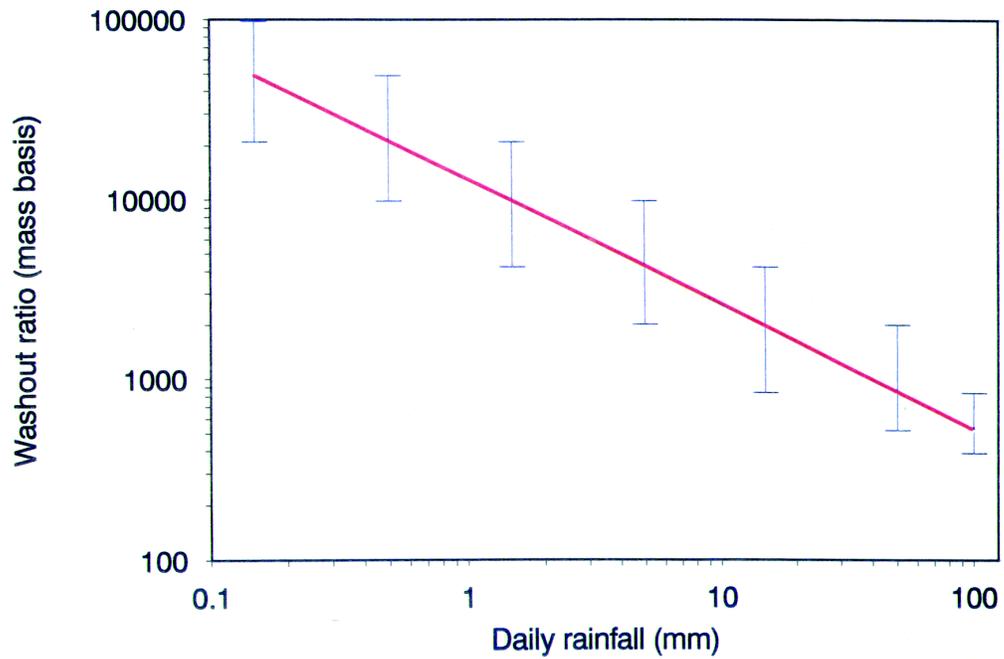
- as a function of distance from the NTS (in the absence of rain) in Figure A7.5,
- as a function of daily rainfall in Figure A7.6.

A7.5. RESULTS

Vegetation deposition densities have been calculated using equations A7.1 to A7.4. The calculations have been made in a static manner, using the distributions indicated previously for each parameter. All parameters have been assumed to be independent, with the exception of the dry deposition velocity, v_g , and the mass interception factor, F_{dry}^* , for iodine associated with particles at the closest distance from the NTS ($X = 10 \text{ km}$). The distributions of the vegetation deposition densities have been calculated for two sets of physico-chemical forms of ^{131}I : (a) uniquely associated with particles, and (b) distributed among particulate, elemental, and organic forms as shown in Table A7.1. Table A7.7 presents the estimated median values as well as the 5 and 95 percentiles, of the vegetation deposition density, A_p in nCi m^{-2} , corresponding to a unit time-integrated concentration in air of 1 nCi d m^{-3} . The values estimated for particles attached to particles and for the assumed mixture of physico-chemical forms, are shown in Figures A7.7 and A7.8 for distances of 100 and 3000 km downwind from the NTS, respectively.

The medians of the ratios of the vegetation deposition densities obtained when ^{131}I is distributed among the three physico-chemical forms and when ^{131}I is attached to particles ($\langle A_p(\text{mix}) / A_p(\text{P}) \rangle$), are presented in Table A7.7 along with the 5 and 95 percentiles of the distributions. The ratios obtained at distances of 100 and 3000 km downwind from the NTS are illustrated in Figure A7.9.

7.3. Variation of the washout ratio as a function of daily rainfall for ^{131}I associated with particles, for a distance of 100 km from the NTS. (The vertical bars represent the assumed ranges of uncertainty)



7.4. Variation of the washout ratio as a function of daily rainfall for ^{131}I in molecular and organic form . (The vertical bars represent the assumed ranges of uncertainty)

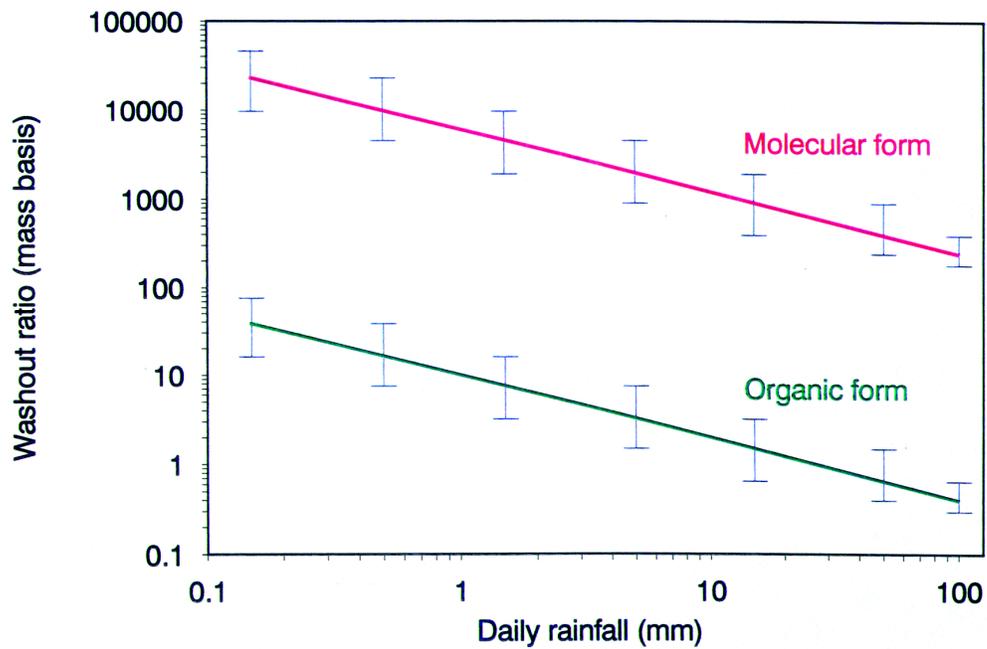


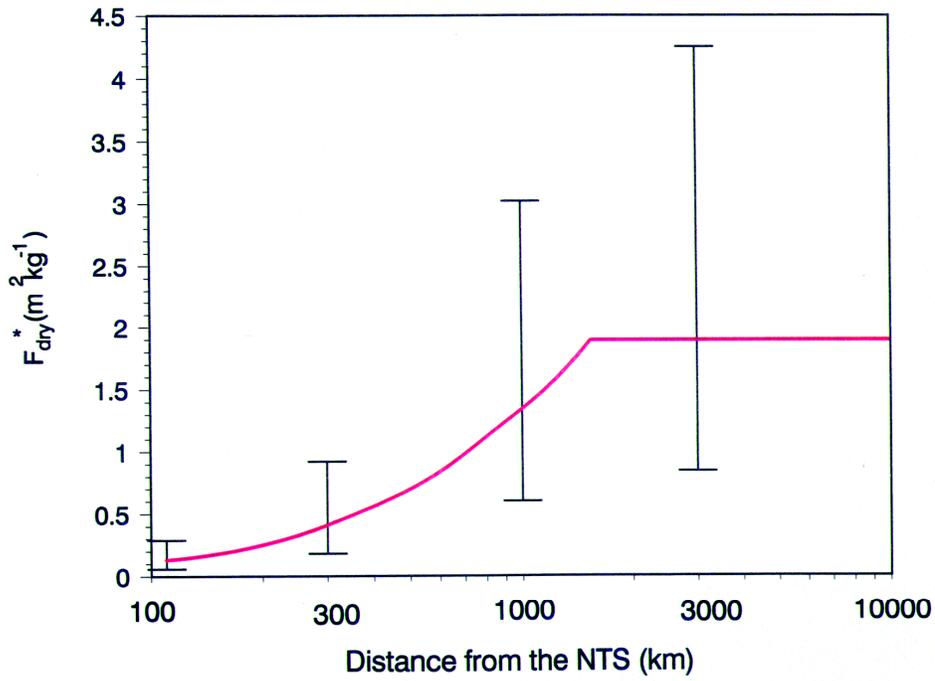
Table A7.5. Best estimates and ranges of wash-out ratios. $WR \{(nCi\ kg^{-1}) / (nCi\ (kg^{-1}))^{-1}\}$ for ^{131}I in the 3 physico-chemical forms and the 4 distances from the NTS, X (km), considered.

| Prec. index | Daily rain mm | Particulate | | | Molecular | | | Organic | | |
|-------------|------------------|-------------|-------|-------|-----------|-------|-------|---------|------|------|
| | | Min. | Mode | Max. | Min. | Mode | Max. | Min. | Mode | Max |
| X = 100 km | | | | | | | | | | |
| 2 | 0.15 | 21000 | 49000 | 98000 | 9700 | 23000 | 46000 | 16 | 38 | 76 |
| 3 | 0.5 | 9800 | 21000 | 49000 | 4500 | 9700 | 23000 | 7.5 | 16 | 38 |
| 4 | 1.5 | 4200 | 9800 | 21000 | 1900 | 4500 | 9700 | 3.2 | 7.5 | 16 |
| 5 | 5 | 2000 | 4200 | 9800 | 900 | 1900 | 4500 | 1.5 | 3.2 | 7.5 |
| 6 | 15 | 840 | 2000 | 4200 | 390 | 900 | 1900 | 0.65 | 1.5 | 3.2 |
| 7 | 50 | 520 | 840 | 2000 | 240 | 390 | 900 | 0.40 | 0.65 | 1.5 |
| 8 | 100 | 390 | 520 | 840 | 180 | 240 | 390 | 0.30 | 0.40 | 0.65 |
| 9 | 150 | 260 | 390 | 520 | 120 | 180 | 240 | 0.20 | 0.30 | 0.40 |
| X = 300 km | | | | | | | | | | |
| 2 | 0.15 | 13000 | 31000 | 62000 | 9700 | 23000 | 46000 | 16 | 38 | 76 |
| 3 | 0.5 | 6100 | 13000 | 31000 | 4500 | 9700 | 23000 | 7.5 | 16 | 38 |
| 4 | 1.5 | 2300 | 6100 | 13000 | 1900 | 4500 | 9700 | 3.2 | 7.5 | 16 |
| 5 | 5 | 1200 | 2600 | 6100 | 900 | 1900 | 4500 | 1.5 | 3.2 | 7.5 |
| 6 | 15 | 520 | 1200 | 2600 | 390 | 900 | 1900 | .065 | 1.5 | 3.2 |
| 7 | 50 | 320 | 520 | 1200 | 240 | 390 | 900 | 0.40 | 0.65 | 1.5 |
| 8 | 100 | 240 | 320 | 520 | 180 | 240 | 390 | 0.30 | 0.40 | 0.65 |
| 9 | 150 | 160 | 240 | 320 | 120 | 180 | 240 | 0.20 | 0.30 | 0.40 |
| X = 1000 km | | | | | | | | | | |
| 2 | 0.15 | 7800 | 18000 | 36000 | 9700 | 23000 | 46000 | 16 | 38 | 76 |
| 3 | 0.5 | 3600 | 7800 | 18000 | 4500 | 9700 | 23000 | 7.5 | 16 | 38 |
| 4 | 1.5 | 1600 | 3600 | 7800 | 1900 | 4500 | 9700 | 3.2 | 7.5 | 16 |
| 5 | 5 | 720 | 1600 | 3600 | 900 | 1900 | 4500 | 1.5 | 3.2 | 7.5 |
| 6 | 15 | 310 | 720 | 1600 | 390 | 900 | 1900 | 0.65 | 1.5 | 3.2 |
| 7 | 50 | 190 | 310 | 720 | 240 | 390 | 900 | 0.40 | 0.65 | 1.5 |
| 8 | 100 | 140 | 190 | 310 | 180 | 240 | 390 | 0.30 | 0.40 | 0.65 |
| 9 | 150 | 90 | 140 | 190 | 120 | 180 | 240 | 0.20 | 0.30 | 0.40 |
| X = 3000 km | | | | | | | | | | |
| 2 | 0.15 | 4900 | 11000 | 22000 | 9700 | 23000 | 46000 | 16 | 38 | 76 |
| 3 | 0.5 | 2300 | 4900 | 11000 | 4500 | 9700 | 23000 | 7.5 | 16 | 38 |
| 4 | 1.5 | 980 | 2300 | 4900 | 1900 | 4500 | 9700 | 3.2 | 7.5 | 16 |
| 5 | 5 | 450 | 980 | 2300 | 900 | 1900 | 4500 | 1.5 | 3.2 | 7.5 |
| 6 | 15 | 190 | 450 | 980 | 390 | 900 | 1900 | 0.65 | 1.5 | 3.2 |
| 7 | 50 | 120 | 190 | 450 | 240 | 390 | 900 | 0.40 | 0.65 | 1.5 |
| 8 | 100 | 90 | 120 | 190 | 180 | 240 | 390 | 0.30 | 0.40 | 0.65 |
| 9 | 150 | 60 | 90 | 120 | 120 | 180 | 240 | 0.20 | 0.30 | 0.40 |

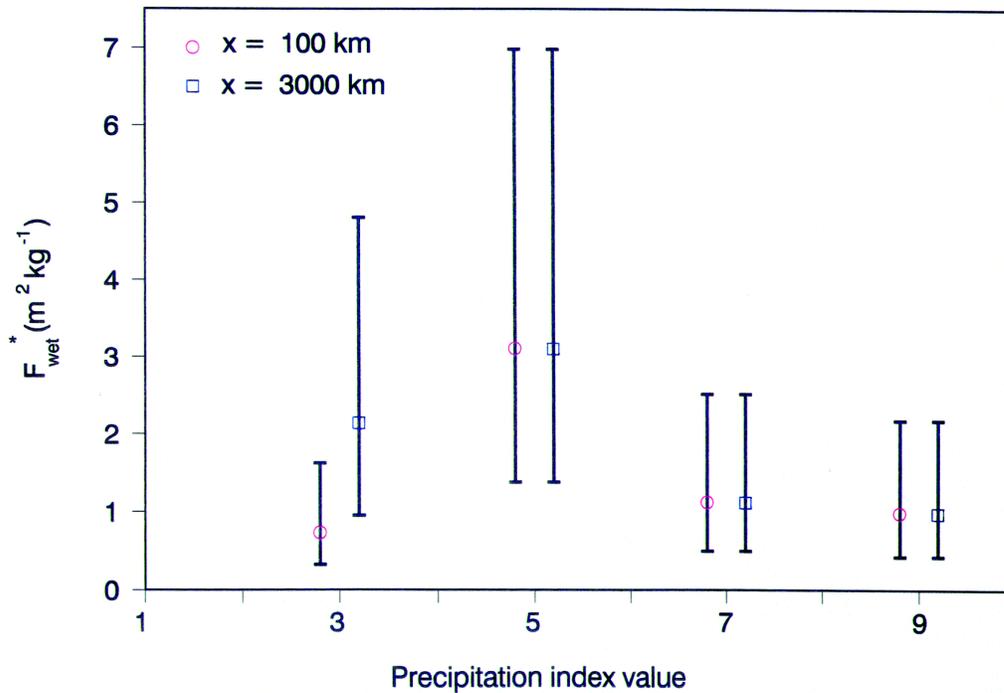
.6. Best estimates of the mass interception factors, F_{dry}^* and F_{wet}^* , in $m^2 kg^{-1}$ (dry).

| Chemical form | Precip. Index value | Distance from the NTS (km) | | | |
|---------------|----------------------|----------------------------|------|------|------|
| | | 100 | 300 | 1000 | 3000 |
| | | $F_{dry}^* (m^2 kg^{-1})$ | | | |
| Particles | 1 | 0.13 | 0.41 | 1.34 | 1.89 |
| Molecular | 1 | | | | |
| Organic | 1 | | | | |
| | | $F_{wet}^* (m^2 kg^{-1})$ | | | |
| Particles | 2 | 0.30 | 0.57 | 1.45 | 1.97 |
| | 3 | 0.72 | 0.95 | 1.70 | 2.14 |
| | 4 | 1.91 | 2.03 | 2.40 | 2.62 |
| | 5 | 3.10 | 3.10 | 3.10 | 3.10 |
| | 6 | 1.63 | 1.63 | 1.63 | 1.63 |
| | 7 | 1.12 | 1.12 | 1.12 | 1.12 |
| | 8 | 1.01 | 1.01 | 1.01 | 1.01 |
| | 9 | 0.97 | 0.97 | 0.97 | 0.97 |
| | Molecular or Organic | 2 | 0.20 | 0.20 | 0.20 |
| 3 | | 0.21 | 0.21 | 0.21 | 0.21 |
| 4 | | 0.26 | 0.26 | 0.26 | 0.26 |
| 5 | | 0.31 | 0.31 | 0.31 | 0.31 |
| 6 | | 0.16 | 0.16 | 0.16 | 0.16 |
| 7 | | 0.11 | 0.11 | 0.11 | 0.11 |
| 8 | | 0.10 | 0.10 | 0.10 | 0.10 |
| 9 | | 0.10 | 0.10 | 0.10 | 0.10 |

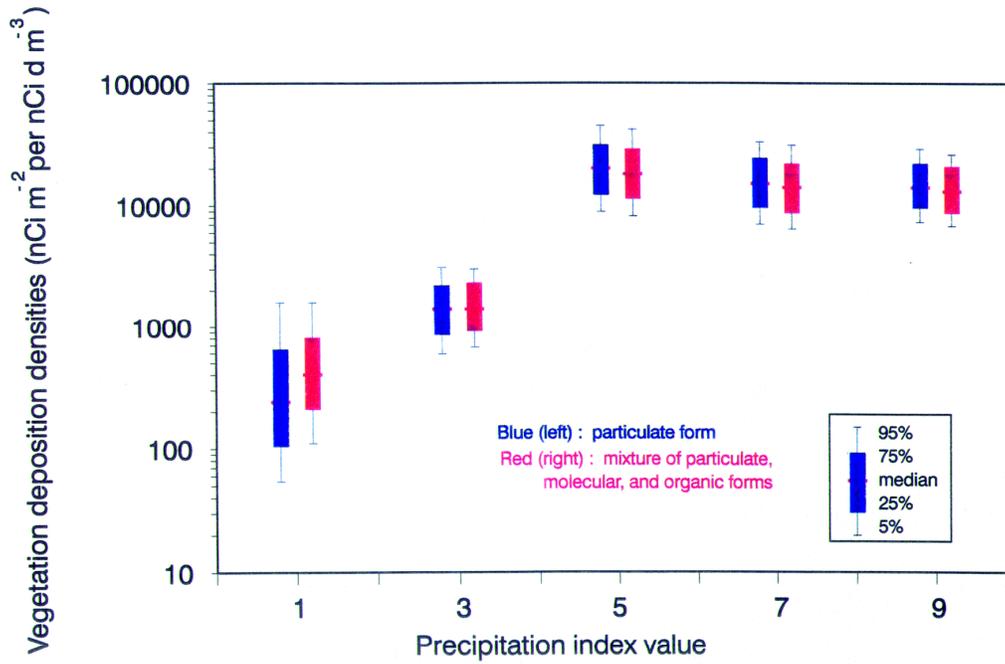
7.5. Variation of the mass interception factor (dry conditions), in $m^2 kg^{-1}$, as a function of distance from the NTS for ^{131}I associated with particles (the vertical bars represent 95% confidence intervals).



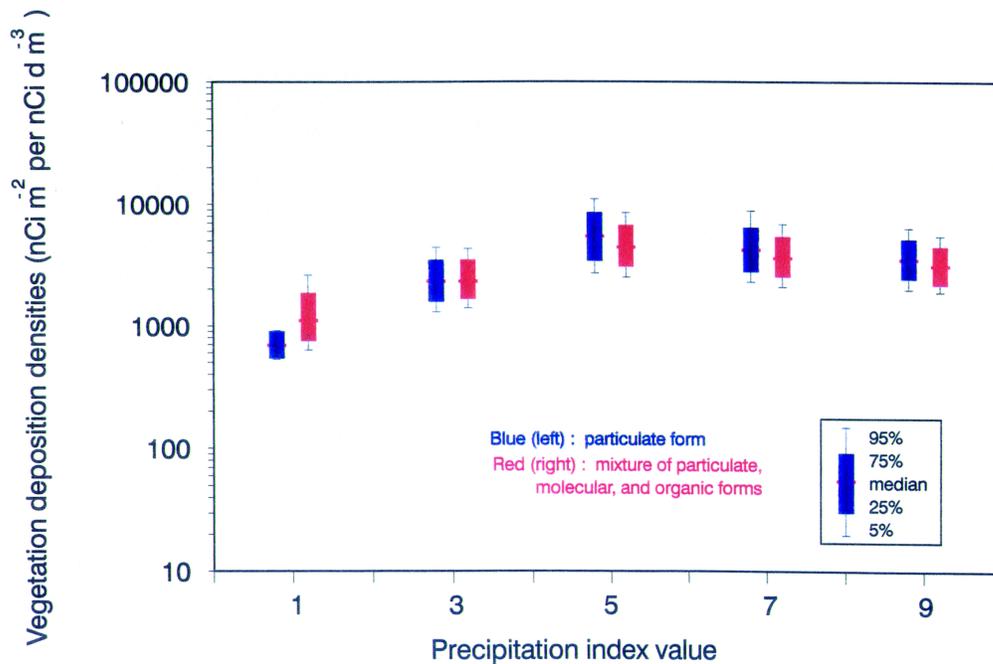
7.6. Variation of the mass interception factor according to the precipitation index value for ^{131}I attached to particles and 2 distances from the NTS. (Median and 95% confidence intervals are shown on the vertical bars)



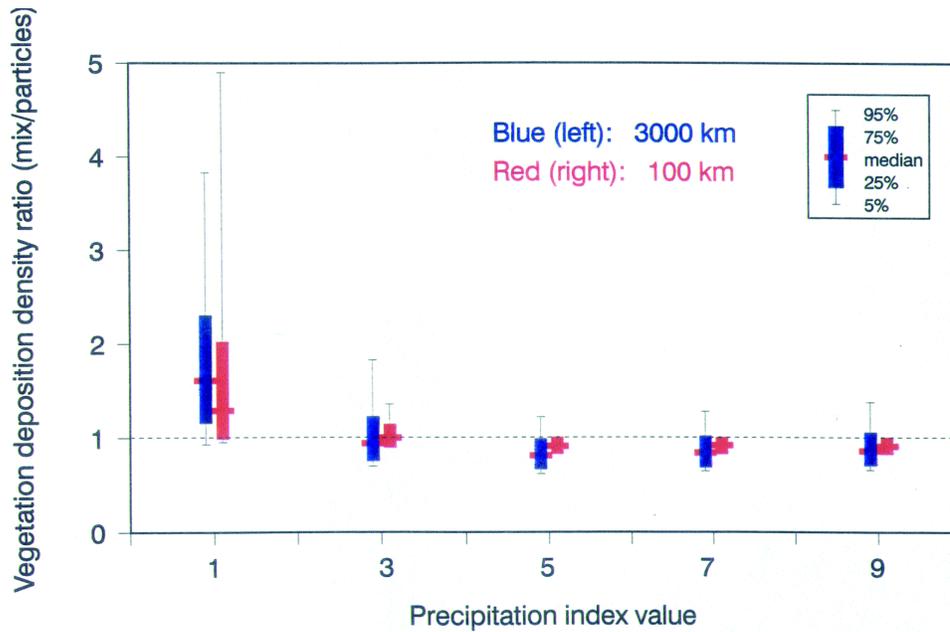
7.7. Variation of the vegetation deposition density as a function of the precipitation index value at the distance of 100 km from the NTS for ^{131}I attached to particles and distributed among 3 physico-chemical forms. Medians, as well as 50% and 95% confidence intervals, are shown on the vertical bars.



7.8. Variation of the vegetation deposition density as a function of the precipitation index value at the distance of 3000 km from the NTS for ^{131}I attached to particles and distributed among 3 physico-chemical forms. Medians, as well as 50% and 95% confidence intervals, are shown on the vertical bars.



7.9. Ratios of normalized vegetation deposition densities (mix/particles) as a function of the precipitation index value for downwind distances from the N of 100 and 3000 km. Medians, as well as 50% and 95% confidence intervals, are shown on the vertical bars.



DISCUSSION

Results of calculations taking into account the distribution of ^{131}I among several physico-chemical forms show that these forms do not produce vegetation deposition densities, therefore, dose estimates that are substantially different from those calculated using the assumption that all ^{131}I was in particulate form. The median values of the ratios of the ^{131}I vegetation deposition densities obtained when ^{131}I is distributed among various physico-chemical forms and when ^{131}I is only in particulate form are in the range of 1.3 to 1.6 for dry deposition and in the range of 0.8 to 1.0 in the presence of precipitation for downwind distances from the NTS. The assumption that is used for these calculations is that ^{131}I is only associated with particles leads, on average, to an underestimation of the thyroid doses from the deposition of ^{131}I occurs in the absence of rain and to a overestimation of the thyroid doses when the deposition of ^{131}I occurs in the presence of rain. However, the uncertainty due to the assumption that ^{131}I is only associated with particles is less than the uncertainties related to the estimation of the number of ^{131}I per unit area of ground. The results shown in Table A7.7 were used to assess the effect of the assumption that all of the ^{131}I was associated with particulate material on dose estimates for locations relatively near and at greater distances. Distributions of daily rainfalls for the period 1951-1960 were extracted from the climatic data for Las Vegas NV, Boise ID, Denver CO, Memphis TN, Jacksonville FL, and Albany NY¹. The median ratios of $[A_p(\text{mix})/A_p(\text{P})]$ in Table A7.7. were used with the distributions of

daily rainfall to estimate weighted biases for these particulate forms. Some linear interpolations and extrapolations of the ratios in Table A7.7. were used to make the estimates because, except for Denver, the distance from the NTS of the location considered differed from the distances for which $[A_p(\text{mix})/A_p(\text{P})]$ had been estimated.

Table A7.8. Median estimates of bias derived from precipitation frequencies during 1951-1960

| Location | Distance from NTS (km) | Estimate of Bias |
|------------------|------------------------|------------------|
| Las Vegas, NV | 150 | 1.3 |
| Boise, ID | 720 | 1.3 |
| Denver, CO | 1020 | 1.1 |
| Memphis, TN | 2330 | 1.3 |
| Jacksonville, FL | 3136 | 1.3 |
| Albany, NY | 3780 | 1.3 |

Table A7.7. Best estimates (medians) and ranges (5 and 95 percentiles) of vegetation deposition densities, A_p [(nCi m⁻²) (nCi d m⁻³)⁻¹], for ¹³¹I associated with particles, A_p (P), and for ¹³¹I distributed according to the assumed mixture of physico-chemical forms, A_p (mix), for a range of daily rainfalls and distances from the NTS. The ratios, A_p (mix)/ A_p (P), also are presented.

| Reference daily rainfall | A_p (P) | | | A_p (mix) | | | A_p (mix)/ A_p (P) | | |
|--------------------------|-----------|--------|-------|-------------|--------|-------|------------------------|--------|------|
| | 5% | Median | 95% | 5% | Median | 95% | 5% | Median | 95% |
| X= 100 km | | | | | | | | | |
| 0 | 54 | 240 | 1600 | 110 | 40 | 1600 | 0.95 | 1.29 | 4.90 |
| 0.15 | 250 | 560 | 1700 | 320 | 670 | 1800 | 0.94 | 1.10 | 2.02 |
| 0.5 | 590 | 1400 | 3100 | 670 | 1400 | 3000 | 0.91 | 1.00 | 1.36 |
| 1.5 | 3700 | 8400 | 19000 | 3500 | 7700 | 17000 | 0.86 | 0.93 | 0.99 |
| 5 | 8900 | 20000 | 45000 | 8200 | 18000 | 42000 | 0.85 | 0.91 | 0.97 |
| 15 | 6300 | 14000 | 22000 | 5800 | 13000 | 29000 | 0.85 | 0.92 | 0.98 |
| 50 | 7000 | 15000 | 33000 | 6400 | 14000 | 31000 | 0.85 | 0.92 | 0.98 |
| 100 | 7500 | 15000 | 31000 | 7000 | 13000 | 28000 | 0.85 | 0.91 | 0.98 |
| 150 | 7300 | 14000 | 29000 | 6800 | 13000 | 26000 | 0.85 | 0.91 | 0.98 |
| X = 300 km | | | | | | | | | |
| 0 | 170 | 410 | 1200 | 330 | 700 | 1700 | 0.95 | 1.51 | 4.22 |
| 0.15 | 630 | 200 | 2400 | 780 | 1400 | 2500 | 0.88 | 1.06 | 1.77 |
| 0.5 | 1200 | 2400 | 5000 | 1300 | 2400 | 4400 | 0.83 | 0.95 | 1.34 |
| 1.5 | 2700 | 5800 | 12000 | 2500 | 5100 | 10000 | 0.79 | 0.88 | 1.04 |
| 5 | 5600 | 13000 | 30000 | 4900 | 11000 | 24000 | 0.77 | 0.84 | 0.92 |
| 15 | 4200 | 9100 | 20000 | 3800 | 7700 | 17000 | 0.78 | 0.86 | 0.97 |
| 50 | 4400 | 9500 | 21000 | 4000 | 8200 | 17000 | 0.78 | 0.85 | 0.95 |
| 100 | 4900 | 9600 | 19000 | 4400 | 8200 | 15000 | 0.78 | 0.85 | 0.94 |
| 150 | 4800 | 9300 | 18000 | 4300 | 7900 | 15000 | 0.78 | 0.85 | 0.95 |
| X = 1000 km | | | | | | | | | |
| 0 | 460 | 760 | 1300 | 590 | 1000 | 2000 | 0.87 | 1.30 | 2.64 |
| 0.15 | 1200 | 1900 | 3300 | 1200 | 1900 | 3300 | 0.78 | 0.96 | 1.57 |
| 0.5 | 1600 | 2800 | 5500 | 1500 | 2600 | 4600 | 0.75 | 0.89 | 1.33 |
| 1.5 | 2300 | 4600 | 9500 | 2200 | 3900 | 7400 | 0.72 | 0.83 | 1.10 |
| 5 | 3800 | 8100 | 18000 | 3300 | 6400 | 13000 | 0.70 | 0.79 | 0.96 |
| 15 | 2900 | 6000 | 13000 | 2600 | 5000 | 9600 | 0.70 | 0.81 | 1.04 |
| 50 | 3300 | 6400 | 13000 | 2900 | 5200 | 10000 | 0.71 | 0.80 | 0.99 |
| 100 | 3500 | 6500 | 13000 | 3100 | 5300 | 9600 | 0.71 | 0.80 | 0.98 |
| 150 | 3400 | 5800 | 11000 | 2900 | 4800 | 8300 | 0.72 | 0.81 | 1.01 |
| X = 3000 km | | | | | | | | | |
| 0 | 530 | 690 | 9100 | 630 | 1100 | 2600 | 0.93 | 1.61 | 3.83 |
| 0.15 | 1000 | 1600 | 2700 | 1100 | 1800 | 3500 | 0.75 | 1.06 | 2.15 |
| 0.5 | 1300 | 2300 | 4400 | 1400 | 2300 | 4300 | 0.70 | 0.94 | 1.83 |
| 1.5 | 1800 | 3300 | 6400 | 1800 | 3000 | 5500 | 0.66 | 0.87 | 1.47 |
| 5 | 2700 | 5400 | 11000 | 2300 | 4400 | 8500 | 0.62 | 0.81 | 1.22 |
| 15 | 2100 | 3900 | 7900 | 2000 | 3400 | 6400 | 0.63 | 0.84 | 1.37 |
| 50 | 2300 | 4200 | 8900 | 2100 | 3600 | 6900 | 0.65 | 0.84 | 1.28 |
| 100 | 2400 | 4000 | 7400 | 2200 | 3500 | 6100 | 0.65 | 0.84 | 1.30 |
| 150 | 2000 | 3500 | 6400 | 1900 | 3100 | 5500 | 0.66 | 0.86 | 1.38 |

gas associated with particulate material may have bias of 10-30% in the dose calculations. Bias estimates of the six locations are shown in *Table A7.8*; the locations listed in order of increasing distance from the NTS. The estimates are generally uniform and do not show a trend on the distance from the NTS. The estimates in the table reflect average precipitation frequency during the 10-year period, not necessarily the conditions under which fallout deposition actually occurred. Detailed results for the important test series, tabulated in the appendix to the report, were used to estimate the mean bias periods when fallout occurred at four locations. These locations were chosen in part because of the relative completeness of records of gummed film deposition results. The estimates of bias are based on the assumption that all the radioiodine was associated with particulates for the four locations are shown in *Table A7.9*.

| 9. Mean estimates of bias derived from records of precipitation during actual fallout depositions from tests at the NTS. | |
|---|---------------------|
| Location | Estimated Mean Bias |
| Denver | 1.3 ± 0.5 |
| Memphis | 1.1 ± 0.2 |
| Detroit | 1.3 ± 0.4 |
| New York City | 1.1 ± 0.2 |

The mean estimate and an approximate standard deviation are given in the table. These more definitive results, based on records of known fallout deposition, indicate that the likely values of the bias are small. The uncertainties in the estimates are not sufficiently large that one cannot conclude that there is no bias at these locations due to the assumption that all the iodine was in particulate form.

This section was kindly performed for the project by Milton Smith (NOAA/Air Resources Laboratory).

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